

Let

$H$  = "height" of atmosphere

$L$  = Pathlength of sunlight through atmosphere.

$\Theta$  = Solar Zenith Angle

$$\text{Then } \cos \theta = \frac{H}{L}, \quad L = \frac{H}{\cos \theta}$$

Let

$m = 1/\cos \theta = \text{Air Mass}, \quad m \geq 1$

$m = 1$  Sun overhead

$m = 10$  Sunrise or sunset

$I_0 = I_0(\lambda) = \text{T.O.A Solar irradiance}$

$I(\lambda) = \text{GND Solar irradiance}$

$\tau$  = optical depth - dimensionless quantity  
(when the sun is overhead)

Then

Direct Beam

(no multiple scattering)

$$I(\lambda) = e^{-\frac{\tau(\lambda)}{\cos \theta}} I_0(\lambda)$$

Sunlight at surface

Sunlight at top of atmosphere

$$\text{So } I(\lambda) = e^{-mZ(\lambda)} I_0(\lambda)$$

Idea:

1. Measure  $I(\lambda)$  at the ground at different times, so with different  $m$  values.
2. The sun output  $I_0(\lambda)$  at the top of the atmosphere remains the same.
3. Unknown:  $T(\lambda)$  Solve for it, learn about the atmosphere with it.

Here's how:

$$\tau(\lambda) = \tau_{\text{gas}}(\lambda) + \tau_{\text{aerosol}}(\lambda) + \tau_{\text{cloud}}(\lambda)$$

$$\tau_{\text{gas}}(\lambda) \Rightarrow \tau_{\text{gas}}^{\text{scattering}}(\lambda) + \tau_{\text{gas}}^{\text{absorption}}(\lambda)$$

Rayleigh Scattering  
 (Blue Sky)

Absorption by gases  
 like water vapor, O<sub>2</sub>,  
 Ozone (O<sub>3</sub>),  
 Nitrogen dioxide (NO<sub>2</sub>)

$$= \frac{a}{\lambda^4}$$

$$\tau_{\text{aerosol}}(\lambda) = \tau_{\text{aerosol}}^{\text{scattering}}(\lambda) + \tau_{\text{aerosol}}^{\text{absorption}}(\lambda)$$

Sulfates	Black Carbon
Nitrates	(Soot)
Sea Salt	Dust
$\sim \frac{b}{\lambda^2}$	Organic Carbon
Dust	$\sim \frac{c}{\lambda}$ (UV wavelengths)
Black Carbon	

Example: Rayleigh Scattering by O<sub>2</sub>, N<sub>2</sub>

Blue  $\frac{\tau_{\text{gas}}^{\text{scattering}}(\lambda = 400 \text{ nm})}{\tau_{\text{gas}}^{\text{scattering}}(\lambda = 600 \text{ nm})} = \left( \frac{600}{400} \right)^4 = \frac{81}{16} \approx 5$

Red  $\tau_{\text{gas}}^{\text{scattering}}(\lambda = 600 \text{ nm})$

Conclusion: Scattering in blue by gases is 5 x scattering in Red.

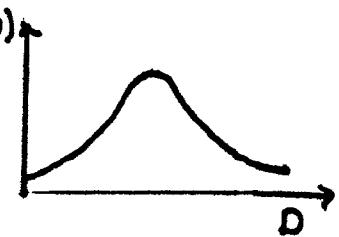
Aerosol:

$$\tau_{\text{aerosol}}^{\text{ext}}(\lambda) = \int_0^H \int_0^\infty \sigma_{\text{ext}}(\lambda, D) N(D, z) dz dD$$

$\sigma_{\text{ext}} = \sigma_{\text{abs}} + \sigma_{\text{sca}}$

Particle extinction cross section as a function of wavelength and particle size.

$N(D, z)$  = Particle Size Distribution  
(D = particle diameter)



Gas Absorption:

$$\tau_{\text{gas}}^{\text{absorption}}(\lambda) = \int_0^H C(z) \sigma_{\text{abs}}(z, \lambda) dz$$

# of  
Absorbing Molecules  
Volume

Absorption Cross  
Section per Molecule

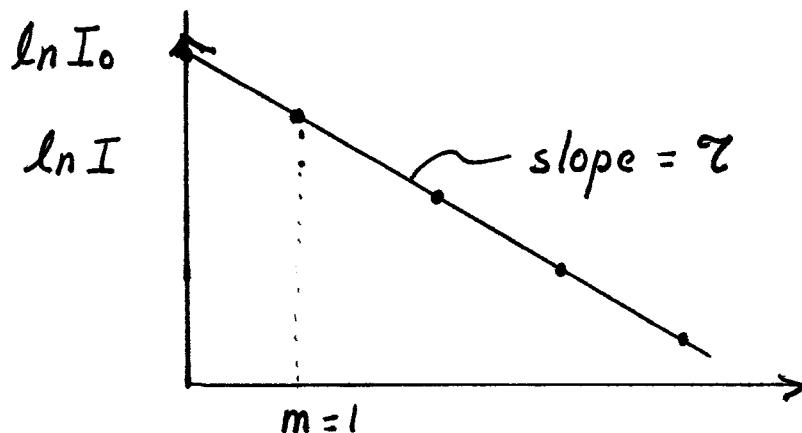
Here's how to crunch the data  
and get information from it.

Date	Time	Lat	Long	Calculate $\theta$	Measure $I(\lambda)$ Volts	Measure $I(\lambda)$ Dark Volts
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$$I(\lambda) = e^{-\tau m} I_0(\lambda)$$

Then  $\ln I = \ln I_0 - m\tau$

offset →                          ↑  
 (fixed,                            independent                          slope use to  
 sunlight at                      variable                            learn about  
 top of                            Atmosphere                       atmosphere



$$m \text{ air mass} = \frac{1}{\cos \theta}$$

## How to process

$$\text{Slope} = \tau = \tau_{\text{gas}}^{\text{sca}} + \tau_{\text{aerosol}} + \cancel{\tau_{\text{cloud}}}$$

look up from      |      }      "clear" day  
 table                aerosol amount  
                     (optical depth)  
 $\tau_{\text{aerosol}} \approx 0.2$  or so.

For 940 nm, also get water vapor concentration

$$\tau = \tau_{\text{gas}}^{\text{scatt.}} + \tau_{\text{H}_2\text{O}}^{\text{absorp.}} + \tau_{\text{aerosol}}$$

|      ↑      ↗  
 look up      Large when      Get from 870 nm  
              atmosphere is moist

Wavelength (nm)	<u>Comment</u>
430	Good for aerosol, Rayleigh scattering larger
470	
530	
660	Good for aerosol, Rayleigh modest
870	Good for aerosol, Rayleigh small
940	water vapor channel

Also, get  $I_o(\lambda)$  from the offset.

$$I_o(\lambda) = k(\lambda) I_o^{\text{calibrated}}(\lambda)$$

↑                          counts or voltage  
calibration factor      from offset

From satellite measurements, TABLE

$$K(\lambda) = \frac{I_o^{\text{calibrated}}(\lambda)}{I_o^{\text{counts}}(\lambda)} \left[ \frac{\text{Watts}}{\text{m}^2 \cdot \text{nm}} \right]$$

Raw output from spectrometer.

Use  $k(\lambda)$  to calibrate any spectrum from the spectrometer.



Light source

